Lecture Notes on Boundary Layer Meteorology

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Abstract

Notes of Lectures and addional information from books: *An introduction to boundary layer meteorology*([1]).

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1. Lecture 1 09/01/2025

1.1 Introduction to Boundary Layer

The Boundary Layer can be defined as part of the troposphere that is directly influenced by the presence of the Earth's surface and responds to surface forcings with a time scale of about an hour or less.



Earth's Surface

Figure 1.1. Atmosphere can be divided into 2 parts: boundary layer near surface and free atmosphere above it.

1.2 Boundary layer forcing mecchanism

What physical process modify boundary layer air parcel?

- 1. Heat transfer to.from the ground.
- 2. Frictional drag.
- 3. Evaporation/transpiration.
- 4. Terrain-induced flow modification.
- 5. Pollution emission.

1.3 Types of air flow or wind

Air flow or wind can be decomposed into following 3 types:

- 1. **Mean Wind** $(\bar{u}, \bar{v}, \bar{w})$: Represents the average wind components in the horizontal (\bar{u}, \bar{v}) and vertical (\bar{w}) directions. It is important for the horizontal transport of quantities such as moisture, heat, momentum, and pollutants, a process known as advection.
- 2. Waves: Atmospheric waves, such as gravity waves, occur mostly at night in the nocturnal boundary layer (NBL). They can influence the structure of the boundary layer and the transport of energy.
- 3. **Turbulence**: The vertical transport of moisture, heat, momentum, and pollutants is primarily dominated by turbulence, which is characterized by chaotic and irregular motion.



Figure 1.2. Plot showing profiles of Mean, Wave and Turbulent winds

1.4 Eddies

Eddies are formed due to the interaction of currents with obstacles like coastlines, underwater topography, or other currents, as well as from the instability of larger current systems. Eddies exhibit a rotational flow pattern, either clockwise or counterclockwise. Eddies can vary from size 100 to 3000 metres and also can exists as small as few millimetres. Small eddies might last for seconds to minutes, while larger oceanic eddies can persist for weeks, months, or even years.

1.5 Turbulence Generation Mechanisms

- **Solar Heating**: Solar heating generates thermals, which are essentially larger eddies that drive turbulence in the atmospheric boundary layer.
- Wind Shear: Variations in wind speed or direction with height create wind shear, which is a significant source of turbulence.
- **Obstacle-Induced Flow**: Deflected flow around obstacles such as trees, buildings, or other structures generates turbulent eddies downstream of these obstacles, creating turbulent wakes.



Figure 1.3. Eddy formation due to Turbulence caused by an obstacle

Large eddies will break into smaller eddies after which small eddies dissipates from K.E. to thermal energy.



Figure 1.4. Variation of Turbulent Kinetic energy with change in Size of eddies

2. Leture 2 15/01/2025

2.1 Taylor's Hypothesis

- When studying atmospheric boundary layer (ABL), It is not easy to create a snapshot of turbulance in the Atmosphere.
- Hence it is easier and cheaper to make measurements of point in the atmoshpere for a longer time, then an instantaneous snapshot.
- So we just consider the atmoshpere is frozen.
- Taylor suggested that turbulence can be considered frozen as it advects past sensor.



Figure 2.1. Eddy propagation





$$\frac{\partial T}{\partial x} = 0.05 \text{K/m}, \quad \frac{\partial T}{\partial t} = -0.5 \text{K/s}$$

$$\frac{DT}{\Delta t} = \underbrace{\frac{\partial T}{\partial t}}_{\text{(Taylor's hypothesis)}} + \underbrace{u \frac{\partial T}{\partial x}}_{\text{(Taylor's hypothesis)}} \quad (2.1)$$

2.2 Virtual Potential Temperature

Virtual potential temperature:

$$\theta_{\nu} = \theta (1 + 0.61r) \tag{2.2}$$

Virtual temperature:

$$T_{\nu} = T(1 + 0.61r) \tag{2.3}$$



Figure 2.3. Pressure v/s Termperature

Question 2.1:

Given 25°C termperature, mixing ratio \bar{r} is 20g/kg, measured Pressurere at 900*hPa*, find virtual potential temperature.

Answer 2.1:

Solution:

$$\theta = T \times \left(\frac{1000}{P}\right)^{0.286} \\ = 298 \times \left(\frac{1000}{900}\right)^{0.286} \\ = 332.222K \\ \theta_{\nu} = \theta \times (1+0.61r) \\ = 332.22 \times (1+0.61 \times 0.025) \\ = 336.273K \\ \theta_{\nu} - \theta \approx 4.05K$$

2.3 Boundary Layer Depth and Structure

- Mixed layer
- Residual layer
- Stable Boundary layer

Capping InversionNauturnal Boundary layer



Figure 2.4. Height vs Local Time Diagram of Atmospheric Boundary Layers.

3. Lecture 3 16/01/2025

Ocean: Variations are minimal, with only 10% changes observed over 1000 km. Significant variations occur primarily during weather phenomena.

Land: Day-to-day and diurnal variations are prominent, with distinct boundary layer structures:

- 1. **Convective Mixed Layer:** Thermodynamically unstable with intense vertical mixing.
- 2. **Residual Layer:** Neutral stratification with turbulence of equal intensity in all directions.
- 3. Stable Boundary Layer (Nocturnal B.L.):
 - Neutral stratification with nocturnal jets (30 m/s, 200 m width).
 - Sporadic turbulence and internal gravity waves transporting air parcels vertically.
- 4. **Capping Inversion:** Found at altitudes between 1.5–3 km, acting as a barrier to upward mixing.
- Entrainment Zone: Transition region from stable to unstable conditions, facilitating energy and mass exchange.

3.1 Stability and Plume Behaviour

Looping plumes: Occur in highly unstable conditions, usually during the day, when strong surface heating causes warm air to rise and interact turbulently with cooler air. This leads to an oscillatory motion that disperses pollutants in multiple directions, often seen in arid regions.

FIGURE for each to be added later.

Fanning plumes: Form in stable conditions, typically at night, when surface cooling creates temperature inversions. The plume spreads horizontally, concentrating pollutants close to the surface, which can impact air quality in urban or industrial areas.

- 3. **Coning plumes:** Develop under neutral conditions, where vertical and horizontal mixing is balanced. The plume takes a cone-like shape, often observed on overcast days or in the early morning and evening, with moderate turbulence.
- 4. **Lofting plumes:** Occur when the atmosphere is stable near the ground but unstable above. Pollutants rise and disperse above the stable layer, reducing ground-level pollution and minimizing surface concentrations.
- 5. **Fumigation plumes:** Happen when pollutants are trapped in a stable layer and then forced downward due to rising turbulence. This leads to high concentrations at the surface, posing risks to air quality, especially in industrial areas.

3.2 Importance of Boundary Layer

The boundary layer plays a critical role in regulating interactions between the Earth's surface and the atmosphere. Its study is important in various fields, including:

- 1. **Agricultural meteorology:** Understanding microclimates within the boundary layer aids in crop management, irrigation planning, and predicting the effects of extreme weather on agriculture.
- 2. Air pollution meteorology: Dispersion and concentration of pollutants are governed by boundary layer processes, making it crucial for air quality monitoring and pollution control strategies.
- 3. **Cloud nuclei meteorology:** The boundary layer provides a reservoir of aerosols and moisture that act as cloud condensation nuclei, influencing cloud formation, precipitation, and local weather patterns.
- 4. **Thunderstorms and hurricanes physics:** The exchange of heat, moisture, and momentum in the boundary layer drives the development and intensity of thunderstorms and hurricanes, making it essential for improving weather prediction models.
- 5. **Urban meteorology:** The boundary layer's interactions with urban landscapes affect local climate, energy balance, and pollutant dispersion, aiding in city planning and sustainability efforts.
- 6. **Renewable energy:** Wind energy potential and efficiency are heavily dependent on boundary layer dynamics, which dictate wind speed profiles and turbulence levels near the surface.

Property	Boundary Layer	Free Atmosphere
Turbulence	Almost continuously	Sporadic, CAT, turbu-
	turbulent	lence within clouds
Friction	Strong drag due to	Small viscous dissipa-
	surface interaction	tion
Dispersion	Rapid turbulent mix-	Small molecular dif-
	ing	fusion
Winds	Near logarithmic pro-	Geostrophic winds
	file	
Vertical	Turbulent vertical mo-	Horizontal transport
Transport	tion	by mean wind
Thickness	100m - 3km (vari-	8-16km (less vari-
	able)	able)

Table 1. Comparison of Boundary Layer and FreeAtmosphere Properties

References

 R. B. Stull. An Introduction to Boundary Layer Meteorology. Atmospheric Sciences Library. Kluwer Academic Publishers, 1988.